



Topography and Depth Influences on Soil Nitrogen and Carbon Sequestration

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Abstract

Topography is one of the important factors affecting soil carbon accumulation in a tropical ecosystem. Slope effects vary in magnitude in different agro-ecosystems. The study was conducted to evaluate the influence of topography on soil nitrogen and carbon sequestration in Mbano Imo State, Nigeria. Three profile pits were dug along a landscape of approximately 200 meters, one on each topographic position. Each profile pit was demarcated into 4 equal depths of 30 cm apart and 3 replicated soil samples were collected from each of the depths for laboratory analyses. Analysis of variance and correlation analysis were used for data analysis. From the results, the average bulk densities and moisture content varied between 1.36-1.59 g cm⁻³ and 6.43-7.27% with the footslope (valley position) containing significantly ($p=0.05$) higher values. 0-30 cm depth had the least bulk density value compared to other depths. At 0-30, 30-60, 60-90 and 90-120 depths, carbon and nitrogen sequestration varied depending on the landscape position. At the depth of 0-30 cm in upslope (summit), mid-slope and footslope, organic carbon, total nitrogen, carbon and nitrogen sequestration were the highest ($p=0.05$). 30-60 cm depth of the footslope sequestered significantly higher carbon (520.30) than that of the summit (376.74 g cm⁻²) and mid slope (26.64) ($p=0.05$). Considering the three different topographic units, the footslope sequestered significantly ($p=0.05$) highest quantities of carbon (629.42 g cm⁻²) and nitrogen (33.18 g Nm⁻²) than the up slope (249.36 g cm⁻², 23.39 g Nm⁻²) and mid slope (170.72 g cm⁻², 14.80 g cm⁻²). Also, organic carbon and total nitrogen were highest in the foot slope compared to the mid and up slopes. Generally, ANOVA result revealed higher availability and sequestration of carbon and nitrogen at 0-30 cm depth and valley position.

Keywords:

Carbon; Nitrogen; Sequestration; Soil Depth; Topography

Introduction

Important factors that control soil carbon levels include topography, climate, hydrology, biological activity, vegetation patterns and land use [1]. However, even within the same ecosystem, soil carbon and nitrogen sequestration potentials differ in response to many other factors such as slope position (topography), management practices and climatic conditions. Soil organic carbon reacts quickly to

anthropogenic activities such as environmental degradation, biomass burning, deforestation and land use changes; and rapid decline in carbon and nitrogen contents of arable soils globally due to these activities has brought about a renewed focus on soil carbon and nitrogen sequestration studies.

Topography influences soil nitrogen and carbon through erosion and redistribution of soil materials through leaching, infiltration and runoff potentials [2,3]. However, the effects of topography on soil carbon are likely to vary in magnitude under agricultural systems with different management practices [2] and soil depth. Soil depth controls soil nitrogen and carbon dynamics by bioturbation, placement of plant and animal residues on the surface of the soil and/or incorporation of organic materials within the epipedon and endopedon of soil.

Soil is an important carbon pool and the most vital component of the biosphere [4]. Soils are regarded as the largest carbon pools of the earth carbon cycle and contained about thrice more carbon than the vegetation. Globally, the topsoil (0-30 cm depth of soil) contains about 1500 pg of carbon and this fraction of carbon is lost annually through soil erosion and other human activities. However, the quantity of CO₂ in the atmosphere increases steadily due to man activities that encourage carbon emissions.

The knowledge of the impact of topography on soil carbon stocks regionally and globally is imperative because, this stock is not only two times the total amount of CO₂-C in the atmosphere, but it is also easily upset by slope factors and agricultural activities taking place on the soil surface (top soil). More so, restoration and maintenance of soil health through soil carbon and nitrogen management has remained a major challenge for tropical soils. To make this successful, the comprehensive knowledge of the sequestration of carbon and nitrogen in the tropical soils should form an essential pre-requisite in future land resource management programs. Therefore, this study investigated the effects of topography and soil depth on carbon and nitrogen sequestration.

Materials and Methods

Study site

The research was carried out at Isiala Mbano, (50 39N, 70 13E) in 2018. According to climatic information obtained from the Nigerian Meteorological Unit, the area receives about 2000 mm average rainfall, 26°C mean annual temperature and approximately 80% relative humidity [5]. Benin formation (the coastal plain sand) is the lithological material from which the soils of Isiala Mbano are derived from [6]. These soils are highly weathered, strongly acid, and coarse-textured and generally low in fertility [6]. Isiala Mbano is rainforest vegetation covered with a lot of shrubs, grasses and trees such as oil and raffia palm trees, oil bean trees, elephant grasses, orchids, rubber, strangler figs, pitcher plants, which capture atmospheric carbon via photosynthesis. The topography of the studied site is a rolling landscape with typical hill slope of about 200 meters.

Soil sampling and analyses

Site selection was done randomly to represent the range of toposequences present in the area. The topographic (terrain) attribute considered was the slope steepness, consisting of the 0.5-1% (summit or upslope), 9-10% (mid-slope) and $\geq 40\%$ (footslope or valley position). A soil profile was dug and sampled in each of the three

different slope positions with the same environmental conditions and at different depths of 0-30, 30-60, 60-90 and 90-120 cm respectively in August 2018. A total of 36 samples (3 profile pits × 4 sampling depths × 3 samples per depth) together with 36 undisturbed soil samples using core sampler for the evaluation of soil bulk density were collected. Soil samples were air dried at room temperature, crushed, sieved and analyzed in the laboratory. Hydrometer method was used in Soil particle fraction determination [7]. Core method described by Grossman and Reinsch [8] was used for bulk density determination while the gravimetric method was used for the determination of soil moisture content. Soil reaction was determined using 1:2.5 soil: water ratio [9]. Available P was extracted using Bray-P 2 solution [10]. Exchangeable bases were determined by the neutral ammonium acetate (NH₄OAc of pH 7) procedure [9]. Total Exchangeable Acidity (TEA) was determined by the procedure [11]. Wet digestion method was used [12] for organic carbon determination while total nitrogen was determined by Kjeldahl digestion method [13]. Carbon sequestration (g Cm⁻²) was calculated using the method of [14]: (BD (g Cm⁻³) × OC (g kg⁻¹) × depth cm) while Nitrogen sequestration (gNm⁻²) was calculated using the method [15]: (BD (g Cm⁻³) × TN (g kg⁻¹) × depth) where, BD=Bulk Density, OC=Organic Carbon and TN=Total Nitrogen.

Statistical analysis

Analysis of variance (ANOVA) was used to test differences in soil properties, soil carbon and nitrogen sequestration across soils of

different slope positions and depths. For statistically different parameters (p=0.05), means were separated using the Least Significant Difference (LSD). Correlation analysis was conducted to check the relationships among soil variables.

Results and Discussion

Soil physical and chemical properties

The slope and depth influences on soil properties are shown in Figures 1-3 and Table 1. Soils of the upslope had more of sandy texture while the textures of the mid and footslopes varied between sandy loam and sandy clay loam respectively. Across all sites, soil depth affected the movement of clay, with the 60-90, 90-120 cm depths containing significantly (p=0.05) higher clay than the other depths (Table 1). At the summit and mid slope, the clay content of 0-30 cm differed significantly from 30-60 cm depth. Increased clay with depth had been reported [16]. The clay contents of the different slope positions differed significantly (p=0.05) from each other (Figure 1). The mean bulk density and gravimetric moisture content differed significantly (p=0.05) from each other. Soil pH ranged between 4.87 and 5.22. Across sites, soil reaction and total exchangeable acidity were affected by slope position at 5% level of confidence. Statistically, the valley position contained the highest total nitrogen (0.12 g kg⁻¹), gravimetric moisture (7.61) and organic carbon (1.33 g kg⁻¹).

Depth (cm)	Sand	Silt (gkg ⁻¹)	Clay	BD g cm ⁻³	MC (%)	OC (gkg ⁻¹)	TN (gkg ⁻¹)	AVP mgkg ⁻¹	C/N	Ph (H2O)	TEB	TEA cmol +kg ⁻¹	EC EC	BS (%)
Summit														
0-30	810	170	20	0.83	7.56	1.97	0.17	5.69	12	5.1	6.26	2.48	8.74	71.6
30-60	740	110	150	1.61	6.65	0.78	0.07	2.53	11	5.33	2.79	1.84	4.63	60.3
60-90	610	80	310	1.4	5.72	0.36	0.03	2.35	12	5.43	5.66	2.2	7.86	72
90-120	620	90	290	1.61	7.71	0.06	0.01	1.86	6	5.03	5.1	0.01	5.11	99.8
Mean	695	112.5	193	1.36	6.91	0.79	0.07	3.11	10	5.22	4.95	1.63	6.59	75.9
Mid-slope														
0-30	810	80	110	1.44	6.21	0.62	0.05	5.72	12	5.43	5.17	1.01	6.18	83.7
30-60	670	140	190	1.48	6.63	0.06	0.01	3.46	6	4.52	9.33	2.44	11.8	79.3
60-90	590	120	290	1.59	7.62	0.36	0.03	2.04	12	5.13	7.68	2.88	10.6	72.7
90-120	550	120	330	1.57	5.25	0.46	0.04	1.96	12	4.38	10.7	2.84	13.6	79.1
Mean	655	115	230	1.52	6.43	0.38	0.03	3.3	10	4.87	8.23	2.29	10.5	78.7
Footslope														

0-30	790	60	150	1.45	9.16	2.65	0.15	6.82	18	4.92	4.62	2.12	6.74	68.6
30-60	710	140	150	1.77	6.26	0.98	0.09	3.72	11	4.93	3.38	2.6	5.98	56.5
60-90	680	100	220	1.55	6.87	0.96	0.09	2.16	11	4.99	8.09	3.08	11.2	72.4
90-120	590	100	310	1.58	6.8	0.84	0.07	1.19	12	5.67	18.6	3.41	22	84.5
Mean	693	100	208	1.59	7.27	1.36	0.1	3.47	13	5.13	8.66	2.8	11.5	70.5
LSD(0.05)	30	10.06	31.8	0.075	0.34	0.25	0.017	0.597	1	0.121	1.39	0.31	2.95	0

Table 1: Physical and Chemical properties of soil.

OC: Organic Carbon; TN: Total Nitrogen; Avp: Available Phosphorus; C/N: Carbon Nitrogen ratio; TEA: Total Exchangeable Acidity; TEB: Total Exchangeable Bases; ECEC: Effective Cation Exchange Capacity; BS: Base Saturation.

Based on the data analysis results, organic carbon and total nitrogen contents of the soils were significantly ($p=0.05$) affected by depth and slope position. The 0-30 cm depth had significantly higher organic carbon and total nitrogen than the other soil depths. The high moisture content of the toe slope may be a consequence of water flow from the summit to the lower position governed by the topographic position, resulting in higher moisture availability at the toe slope [17]. However, soil properties such as available phosphorus showed no significant difference across the slopes and soil depths. The mean TEB and ECEC of the soils ranged from 4.95-8.66 and 6.59-11.46 $\text{cmol}+\text{kg}^{-1}$. From statistical results, the mid and foot slopes had significant higher ($p=0.05$) total exchangeable bases and ECEC than the summit (Figure 3). C/N ratio of the valley slope differed significantly from the other slope positions.

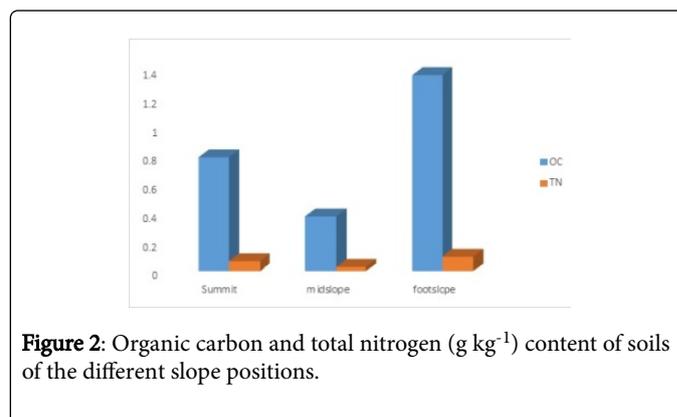


Figure 2: Organic carbon and total nitrogen (g kg^{-1}) content of soils of the different slope positions.

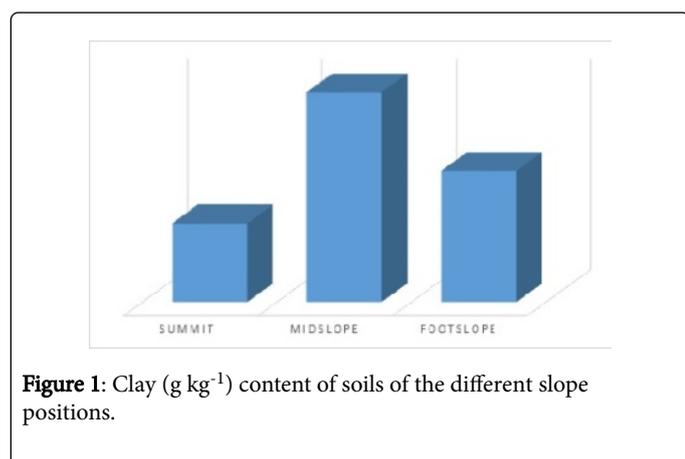


Figure 1: Clay (g kg^{-1}) content of soils of the different slope positions.

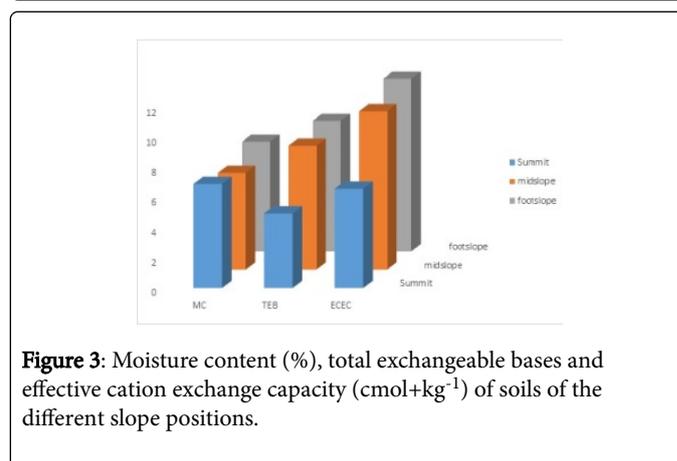


Figure 3: Moisture content (%), total exchangeable bases and effective cation exchange capacity ($\text{cmol}+\text{kg}^{-1}$) of soils of the different slope positions.

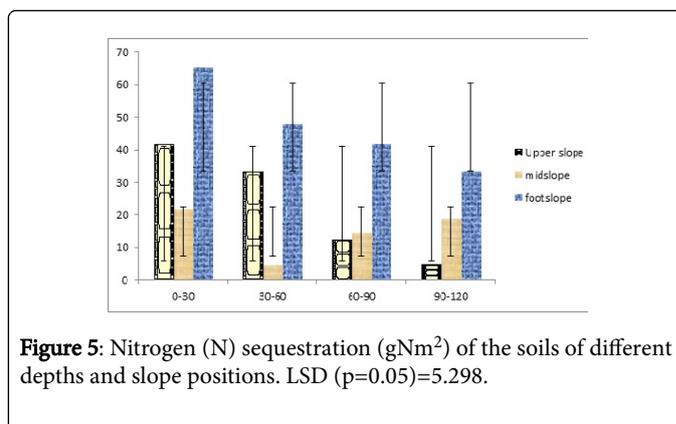
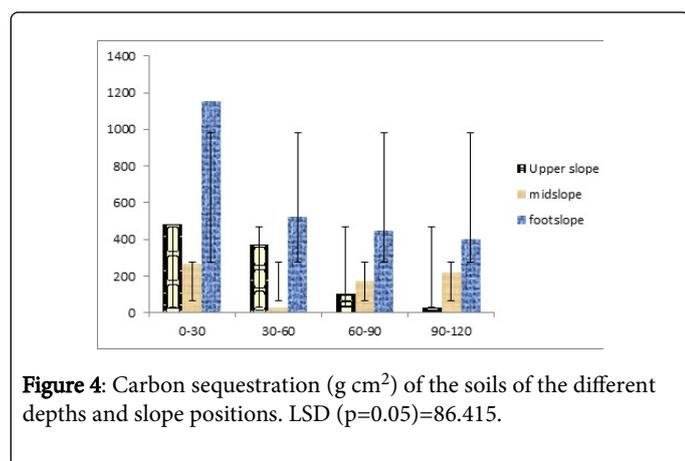
Carbon and nitrogen sequestration

The carbon and nitrogen stocks across the slopes and depths are presented in Figure 4 and Figure 5 respectively. The average carbon sequestered at different topographic positions ranged from 123.4 g Cm^2 at the mid slope to 569.84 g Cm^2 at the valley position. At the summit and valley position, carbon and nitrogen sequestration declined significantly with depth ($p=0.05$), while at the mid slope, the

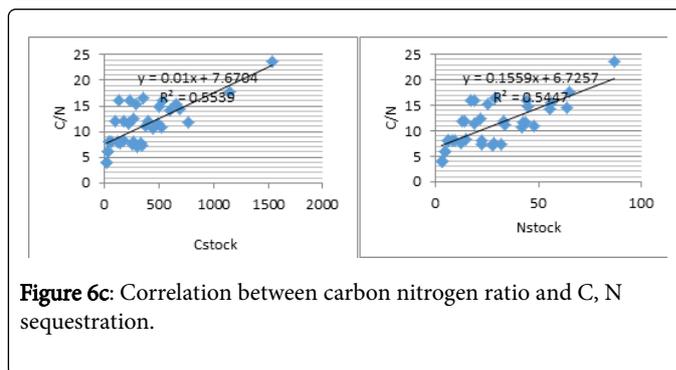
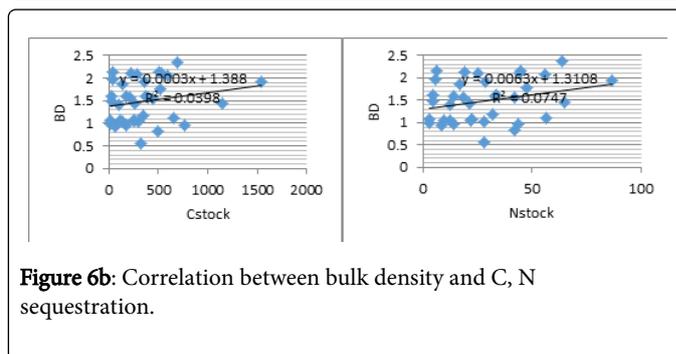
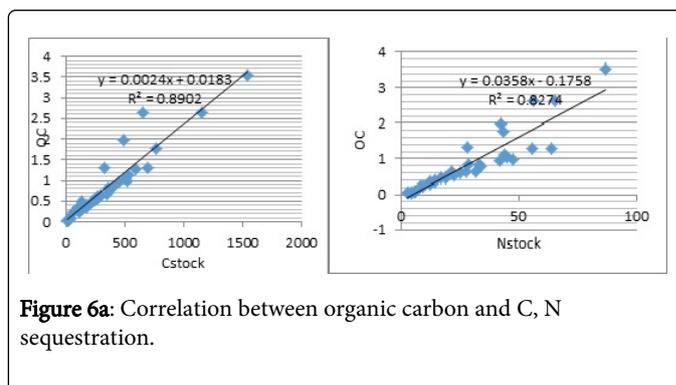
0-30 and 90-120 cm depth sequestered significantly higher carbon. The higher amounts of carbon and nitrogen sequestered in the 0-30 cm depth (top soil) can be empirically explained. The top soil is the first beneficiary of the photosynthetic extraction of carbon into the terrestrial ecosystem from the atmosphere through phyto-mechanisms. Statistically, the slope position affected soil carbon and nitrogen sequestration. The carbon and nitrogen sequestered at the foot slope differed significantly from the summit and mid slope. Significant effect of slope on carbon sequestration has been reported [16,18-20]. Generally, from the results, slope steepness shows to be an effective tool to explain patterns in soil carbon accumulation. Across topographic units, the $\geq 40\%$ slope leads to greater SOC accumulation when associated with relatively moderate slope (9-10%) and flat land ($<1\%$).

High accumulation of carbon at the foot slope could be a consequence of the deposition of organic materials from the upper part of the slope to the valley position by running water. From field observation, the foot slope was densely populated with vegetation than the mid and up slope. The differences in carbon and nitrogen stocks of the soils of the three slope positions are appreciable and may be linked to the heterogeneity of vegetative cover, slope steepness and agronomic activities. The lower carbon and nitrogen sequestered at the summit and mid slope has been reported [21], and could be a consequence of soil conditions that favour rapid decomposition at the valley. At valley positions, carbon accumulation and burial lead to greater soil carbon sequestration and protection of soil organic carbon from further decomposition [22-25]. Some researchers also reported that, at valley positions, the soil carbon that accumulates is not necessarily stabilized on the mineral surfaces; rather, greater carbon stocks at valley positions could be simply protected from microbial decomposition through burial [26,27].

The correlation results are shown in Figures 6a-6e. Statistically, moisture content, available phosphorus, organic carbon and carbon-nitrogen ratio correlated positively with carbon stock ($r=0.532^*$, 0.667^{**} , 0.943^{**} and 0.744^{**}). Also, nitrogen stock had significant positive correlation with total nitrogen ($r=0.906^{**}$), carbon-nitrogen ratio ($r=0.738^{**}$), moisture content and organic carbon (0.924^{**}).



From the results, it could be argued that since moisture is known to influence carbon and nitrogen sequestration, high moisture availability at the foot slope could explain the high carbon and nitrogen sequestration at the foot slope. This was further explained by the significant positive correlation ($r=0.532^*$, 0.531^*) results obtained.



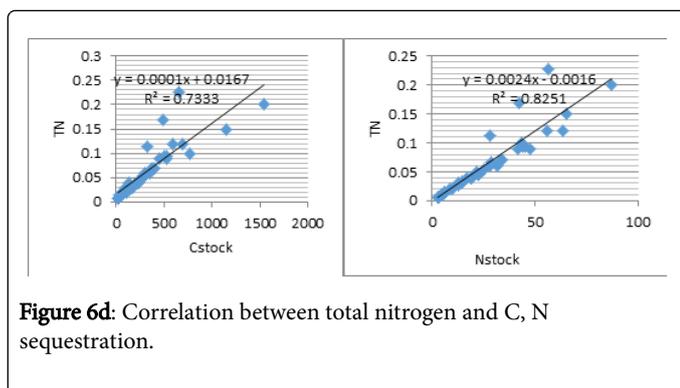


Figure 6d: Correlation between total nitrogen and C, N sequestration.

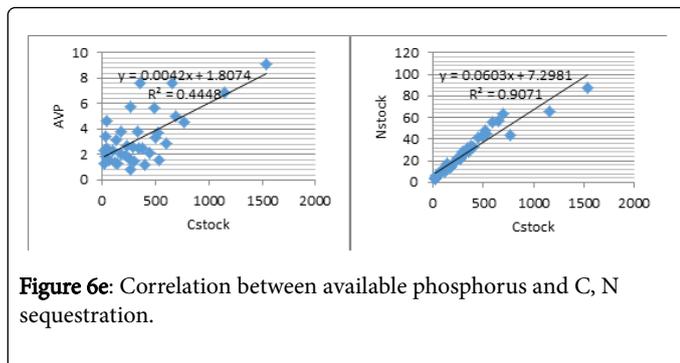


Figure 6e: Correlation between available phosphorus and C, N sequestration.

Conclusion

Studies of soil carbon and nitrogen accumulation in different slope positions can be used as a point of reference for estimating the influence of toposequence on soil carbon and nitrogen sequestration. From our findings, the slope steepness and soil depth affected soil nitrogen and carbon sequestration. The top soil (0-30 cm) and the foot slope sequestered significantly higher quantity of C and N. Generally, our findings support other studies which stated high C and N accumulation at the top soil and valley position. More so, the rate of carbon and nitrogen sequestration at different land units across a toposequence can still be investigated.

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References

- Bhattacharyya T, Pal DK, Mandal C, Velayutham M (2000) Organic carbon stock in Indian soils and their geographical distribution. *National Bureau of Soil Survey and Land Use Planning. Current Science* 79: 655-660.
- Senthilkumar S, Kravchenko AN, Robertson GP (2009) Topography influences Management system effects on total soil carbon and nitrogen. *Soil Sci Soc Am J* 73: 2059-2067.
- Creed IF, Trick CG, Band LE, Morrison IK (2002) Characterizing the spatial heterogeneity of soil carbon and nitrogen pools in the Turkey Lakes Watershed: A comparison of regression techniques. *Water Air Soil Pollution Focus* 2: 81-102.
- Ogle SM, Jay Breidt F, Keith Paustian (2005) Agricultural management impacts on soil organic carbon storage under moist and regions. *Biogeochemistry* 7: 87-121.
- NIMET (Nigerian Meteorological Agency), Nigeria (2016) Climate Weather and Water Information, for sustainable development and safety.
- FDALR [Federal Department of Agricultural Land Resources] (1985) The reconnaissance soil survey of Imo State (1: 250, 000). Owerri: Federal Department of Agricultural Land Resources. p: 133.
- Gee GW, Or D (2002) Particle size distribution. In: Dane J/H and GC Topp (eds) *Methods of soil analysis. Part 4. Physical and mineralogical methods.* Soil science society of America book series, No 5 ASAS and SSA. Madision. pp: 255-293.
- Grossman RB, TG Reinsch (2002) Bulk density. In: Dane, JH and GC Topps (eds) *Methods of soil Analysis. Part 4, Physical methods.* Soil Sci Soc. An Book series No 5, ASA, and SSSA, Madision, WI. pp: 201-228.
- Thomas GW (1996) Soil pH and Soil Acidity. In: *Methods of soil analysis, part 3- chemical methods.* LD Sparks (ed) SSSA Book Series No 5.
- Olsen RS, Sommer LE (1982) Phosphorus in: *methods of soil Analysis part 2* (ed Page, AL, Miller,RH and Keeney, DR) American Soc. Of Agronomy Madision, Wisconsin. pp: 15-72.
- Mclean EC (1982) Aluminum in method of soil analysis. *Agronomy No. 9 part 2 American Society of Agronomy, Madision Wisconsin.methods of soil analysis part 4. Physical methods.* CA Blacks, Dane JH, GC Topp (eds) *Soil Science Soc. Am Book series No 5 ASA and SSA Madision WI.* pp: 201-228.
- Nelson DW, Sommers LE (1996) Total carbon, organic carbon and organic matter. In: *Methods of Soil Analysis. Part 3 Chemical Methods.* Soil Science Society of American Book series. Nsukka. p: 103.
- Bremner JM, Mulvaney CS (1982) Total-Nitrogen. In: *Methods of Soil analysis, Part 2 AL Page, RH Mille and DR Keeney (eds) American Society of Agronomy Madision WI.* pp: 595-624.
- Batjes NH (1996) Total C and N in soils of the world. *Eur J Soil Science* 47: 151-163.
- He Nianpeng, Yunhai Z, Dai J, Han X, Baoyin T, et al. (2012) Land use impact on soil carbon and nitrogen sequestration in typical steppe ecosystems, Inner Mongolia. *J Geogr Sci* 22: 859-873.
- Fissore C, Dalzell BJ, Berhe AA, Voegtle M, EvansM, et al. (2017) Influence of topography on soil organic carbon dynamics in a Southern California grassland. *Catena* 149: 140-149.
- Feras M Ziadat, Awni Y Taimeh, Butros I Hattar (2010) Variation of soil physical properties and moisture content along toposequences in the arid to semiarid area. *Journal Arid Land Research and Management* 24: 81-97.
- Garcia-Ruiz JM (2010) The effects of land uses on soil erosion in Spain: a review. *Catena* 81: 1-11.
- Lasanta T, Arnaez J, Oserin M, Ortigosa LM (2001) Marginal lands and erosion in terraced fields in the Mediterranean Mountains. *Mt Res Dev* 21: 69-76.
- Kazi Kaimul Islam, Somchai Anusontpornperm, Irb Kheoruenromne, Suphicha Thanachit (2018) Carbon sequestration in relation to topographic aspects and land use in Northeast of Thailand. *International Journal of Environment and Climate Change* 8: 138-151.

21. Yoo K, Amundson R, Heimsath AM, Dietrich WE (2006) Spatial patterns of soil Organic carbon on hill slopes: integrating geomorphic processes and the biological C cycle. *Geoderma* 130: 47-65.
22. Doetterl S, Berhe AA, Nadeu E, Wang Z, Sommer M, et al. (2016) Erosion, deposition and soil carbon: a review of process-level controls, experimental tools and models to address C cycling in dynamic landscapes. *Earth Sci Rev* 154: 102-122.
23. Smith SV, Renwick WH, Buddemeier RW, Crossland CJ (2001) Budgets of soil erosion and deposition for sediments and sedimentary organic carbon across the conterminous United States. *Glob Biogeochem Cycles* 15: 697-707.
24. Berhe AA, Kleber M (2013) Erosion, deposition, and the persistence of soil organic matter: Mechanistic considerations and problems with terminology. *Earth Surf Process Land* 38: 908-912.
25. Fontaine S, Barot S, Barre P, Bdioui N, Mary B, et al. (2007) Stability of organic carbon in deep soil layers controlled by fresh carbon supply. *Nature* 450: 277-281.
26. Berhe AA, Harden JW, Torn MS, Kleber M, Burton SD, et al. (2012) Persistence of soil organic matter in eroding versus depositional landform positions. *J Geophys Res and Biogeoscience* 117: 1-61.
27. Berhe AA (2012) Decomposition of organic substrates at eroding vs. depositional landform positions. *Plant Soil* 350: 261-280.